models is the underlying species data used for calibration. Under-representation of distinct environments within a species’ range, or overlooked populations that possess adaptations to local conditions, may lead to reduced predictive power of niche models. Vast numbers of locational records are thus needed to address broad-scale macro-ecological questions with statistical niche models, and a standard practice involves the use of global databases, such as the Global Biodiversity Information Facility (www.gbif.org). Ecosystem monitoring datasets that had not been deposited in this database, such as the ones presented by Ingvaldsen et al., were thus not used to build models for Wisz and colleagues. The resulting paucity of cod occurrence records for the Barents Sea affected the forecast of cod distribution in that area (Fig. 1a,c), especially for EC (Fig. 1a). After updating and refitting the model with Norwegian trawling data from additional, freely available sources of spatially explicit data. Distributions of fish species have already changed in recent years near or at the so-called high-latitude ‘Arctic gateways’ (that is, connections between the Atlantic and Pacific oceans). High-quality distribution data is still not systematically archived in open access biodiversity data portals, and the quality of the available data can be a challenge for many species (for example, ref. 11). Improving the availability of high-quality data in biodiversity data portals will lead to strengthened predictions of distributions, abundance and ecosystems that will serve science and society. Finally, including more occurrences in proximity to the Arctic gateways may unveil a faster rate of interchange than previously estimated, highlighting the need for more systematic transfer of data from ongoing long-term research and monitoring programmes to global databases.

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COMMENTARY:

Playing hide and seek with El Niño

M. J. McPhaden

A much-anticipated ‘monster’ El Niño failed to materialize in 2014, whereas an unforeseen strong El Niño is developing in 2015. El Niño continues to surprise us, despite decades of research into its causes. Natural variations most probably account for recent events, but climate change may also have played a role.

The scientific community has invested considerable effort over the past 50 years in studying El Niño, ever since Jacob Bjerknes first described unusual warm events in the tropical Pacific as the consequence of coupled interactions between the ocean and the overlying atmosphere. El Niño and its cold counterpart La Niña represent the strongest year-to-year climate fluctuation on the planet. What has motivated so much interest in these climatic siblings (which we collectively refer to as the El Niño/Southern Oscillation, or ENSO, cycle) is not only the quest to understand how they work, but also a societal imperative to accurately predict their evolution to help anticipate impacts on lives, property, economic activity and the environment.
Box 1 | Mechanisms for El Niño development.

To understand El Niño, we first need to define what is considered normal in the tropical Pacific. The trade winds, which are a relatively steady flow of air from east to west, drive ocean currents westwards beneath them (see figure). These currents drain warm surface water heated by the sun from the eastern Pacific and pile it up in the western Pacific. In response to this zonal redistribution of upper-ocean water mass, the thermocline (a region of sharp vertical temperature gradient separating the warm, sunlit surface layer from the cold, deep interior ocean) shoals in the east and is pushed down in the west. The shallowness of the thermocline in the eastern Pacific facilitates the upward transport of cold water by the trade winds, a process that is referred to as equatorial upwelling. Upwelling results in a cold tongue of surface water that extends from the west coast of South America to near the International Date Line. The east–west surface temperature contrast reinforces the westward-flowing trade winds, because atmospheric surface pressure is higher over the cooler water in the east, which drives the trade winds westwards.

As the trade winds flow from east to west, they pick up heat and moisture from the ocean. The warm, humid air mass becomes less dense and rises over the western Pacific warm pool, where deep convection leads to towering cumulus clouds and heavy precipitation. Ascending air masses in this region of deep convection return eastwards in the upper levels of the troposphere, then sink over the cooler water of the eastern Pacific.

Every few years, however, the trade winds weaken. This weakening is not a slow and uniform process, but rather one that is punctuated by episodic westerly wind bursts. Wind bursts serve as both a trigger for El Niño onset and an amplifier of ensuing warm SST anomalies.4–6,16,17

They lead to warming along the Equator through two complementary mechanisms. The first is that they force strong eastward-flowing currents that carry warm surface waters from the western Pacific into the central Pacific. The second is that they excite eastward-propagating, downwelling equatorial Kelvin waves that take about 45 days to cross the basin. Kelvin waves focus their energy within a few degrees of the Equator and, in their wake, leave a depressed thermocline over thousands of kilometres in the east–west direction. The depressed thermocline reduces the normal upwelling of cool water to the surface. The net effect of these processes is the development warm SST anomalies along the Equator between the date line and the coast of South America (see figure).

Once warm temperature anomalies are established along the Equator, they can feed back to the atmosphere to cause further weakening of the trade winds11. In the tropics, heated moisture-laden air rises over warm SSTs. As the central Pacific warms up during El Niño, the centre of deep atmospheric convection and heavy rainfall normally located in the far western Pacific migrates eastwards with the warm water. Surface air flow into the convective centre from the west leads to further weakening of the trade winds, which leads to further surface warming. The ocean and the atmosphere thus become locked in a mutually reinforcing feedback loop in which anomalous winds and SSTs continue to amplify. Delayed negative feedbacks involving ocean dynamics eventually take over to terminate the El Niño11,12, returning the system to normal or sometimes causing it to overshoot into a cold La Niña phase. It is the shifting patterns of deep convection and rainfall associated with the anomalous SSTs that lead to the global climatic impacts of El Niño and La Niña. What distinguished 2014 in this scenario was the failure of the positive feedbacks between surface winds and SST to take hold, even though conditions seemed otherwise favourable.

How ENSO may change in the future owing to greenhouse gas forcing is likewise one of the most compelling problems in climate research today.3 Variations in Pacific Ocean circulation associated with ENSO can have profound effects on marine ecosystems and commercially valuable fish stocks. ENSO events also disrupt the circulation of the global atmosphere, increasing the probability of floods, droughts, heat waves and other extreme weather events in far-flung corners of the planet. Sustained and coordinated international research efforts over the past few decades have largely paid off, with scientific progress being measured in terms of new theories for the ENSO cycle, development of sophisticated computer models for seasonal forecasting and establishment of extensive ocean–atmosphere observing systems for tracking ENSO cycle variations. So it was doubly perplexing when, first, an incipient El Niño of major proportions loomed large on the horizon in early 2014, only to disappear suddenly from the radar screen; then, in early 2015, remnants of weak 2014 warming unexpectedly flared up again, instead of quietly fading into history.

What happened? During January to April 2014, a series of strong westerly wind bursts, which are abrupt relaxations of the trade winds lasting typically one to three weeks, occurred in the western equatorial Pacific4,5 (see Box 1). The ocean response to the wind bursts in early 2014 (Fig. 1a) looked strikingly similar to that observed during the onset of the 1997–1998 El Niño (Fig. 1b), which is the strongest on record...
so far\(^5\). The recurrence rate of extreme El Niños such as the 1997–1998 event is once every 15–20 years, leading some experts to interpret this similarity as evidence that we were in store for a big one. Another indicator that conditions were favourable for El Niño in 2014 was the accumulation of excess heat in and above the thermocline along the Equator (Fig. 2). A build up of excess heat content has preceded, by one to three seasons, every moderate to strong El Niño in the past 35 years\(^5\), and the heat content anomaly in early 2014 was higher than any time since 1997. According to El Niño theory\(^4\), this accumulation of excess heat preconditions the ocean to El Niño onset, which can be triggered by episodic westerly wind burst forcing. Therefore, computer models used for seasonal forecasting, especially those that incorporated subsurface temperature data in their initial conditions, predicted the development of El Niño sea surface temperature (SST) warming with a high degree of confidence during the second half of 2014. For example, based on model forecasts from June 2014 initial conditions, the National Oceanic and Atmospheric Administration (NOAA) Climate Prediction Center\(^8\) issued an advisory in early July that indicated, “the chance of El Niño is about 70% during the Northern Hemisphere summer and is close to 80% during the fall and early winter\(^9\).” Given the extraordinary observations and the general agreement among forecasts models, the scientific community and the popular press were abuzz with the possibility that a monster El Niño\(^{10}\) was incubating in the tropical Pacific.

But then the supposedly big El Niño went bust. The trade winds unexpectedly strengthened during the summer and, although westerly wind bursts continued to occur throughout 2014, they were far weaker than in 1997. SST anomalies (that is, deviations from climatological norms) began to decay in the eastern Pacific in mid-2014 in the face of strengthened trades, and seasonal forecasts became increasingly uncertain as the year went on. As a result, headlines changed from bullish to much more subdued and sceptical in heralding...
El Niño’s arrival. By the end of 2014, even though warm conditions prevailed throughout much of the tropical Pacific (Fig. 1c), there was no consensus that an El Niño was underway, given the weakness of the warming and the inconsistency of various key oceanic and atmospheric indicators. Even conceding that the event qualified as an El Niño, the magnitude of the SST anomalies in the season of typical peak development, namely November to January, would rank it among the weakest in the past 60 years (Fig. 3a). Based on year-end observations and model forecasts, NOAA’s January 2015 El Niño advisory stated that weak warm conditions would continue for a couple more months, then die out in early Northern Hemisphere spring.

However, the much-ballyhooed El Niño, though moribund, was not completely dead. Surprisingly, it came roaring back with renewed vigour during the first half of 2015, energized by a new spate of strong westerly wind bursts. These episodic wind events, some of which were associated with highly unusual cyclone activity in the western Pacific in early 2015, are a form of stochastic forcing for El Niño. There is a significant degree of randomness in their timing, amplitude and duration, but they nonetheless preferentially occur over the warmest surface waters of the tropical Pacific. Hence, lingering warm SST anomalies near to and west of the date line in early 2015 may have helped to spark the re-emergence of these wind bursts, leading El Niño to rise, phoenix-like, from the ashes. As a measure of how rapidly El Niño conditions re-intensified in 2015, the +1 °C SST anomaly in June 2015 in the NINO-3.4 index region was the highest for any June since the 1950s, exceeded only by 1997 (Fig. 3b). It is known that ENSO predictions for conditions extending into and through the boreal spring are of relatively low reliability. Even so, the turnaround was remarkable, considering that in early 2015 the event was expected to be over within a couple of months. In contrast, NOAA’s July 2015 advisory note left little room for doubt: “There is a greater than 90% chance that El Niño will continue through Northern Hemisphere winter 2015–2016... with many multi-model averages predicting a strong event...”

Thus, after a faltering start in 2014 and a sudden rebound in 2015, the climate system seems to be heading for an extended two-year period of warming in the tropical Pacific, much like the unusual 1986–1987 El Niño (Fig. 3a). Many El Niño experts were fooled by these developments, both when the widely anticipated monster El Niño went into steep decline and again when it re-ignited with such startling intensity. This puzzling evolution has generated considerable soul-searching about why El Niño forecasts, as well as expectations based on conventional wisdom, fell so wide of the mark.

A key challenge facing the scientific community is to understand why the big event failed to materialize in 2014 when conditions otherwise seemed so favourable. The absence of a sustained feedback between zonal winds and SST in 2014 was symptomatic of the atmosphere’s unexpected insensitivity to warming SSTs in the central equatorial Pacific. As Bjerknes originally pointed out, the wind–SST feedback associated with El Niño is mediated by shifts in deep atmospheric convection and rainfall patterns. However, for the six-month period from June to November 2014, a time when these feedbacks should have promoted growing anomalies in the ocean and atmosphere, persistent deep convection did not take hold in the central Pacific. The atmosphere remained relatively dry, despite anomalously warm SSTs, in sharp contrast to the heavy rains that occurred near to and east of the date line in 1997 (Fig. 1e,f). This lack of strong coupling between the atmosphere and the ocean is the proximate reason for why a monster El Niño did not materialize in 2014.

A complete explanation for what happened must go further though, as we need to address the question of why the Pacific Ocean and overlying atmosphere failed to fully engage with one another. We don’t have to look far for possible suspects that could have conspired to derail the developing El Niño in 2014. There are several possibilities (a few of which will be mentioned here) that may have contributed in varying degrees. The first is the null hypothesis of random chance: while the first few wind bursts energized the initial growth of SST anomalies, the subsequent wind bursts needed to amplify those anomalies failed to develop...
for whatever reason. Another possibility is that the event got underway a month earlier in the calendar year than in 1997, which was perhaps too early, as seasonally varying background conditions favour fastest SST anomaly growth in the second half of the year. In the meantime, negative feedbacks that normally damp El Niño SSTs may have had an opportunity to erode the incipient warm anomalies before they grew too large. Also, for much of 2014, the Indian Ocean Dipole was in a negative phase, mainly due to unusually warm SSTs entrenched off the coast of Java and Sumatra in the eastern equatorial Indian Ocean. These warm SSTs may have helped to anchor deep convection over the Indo-Pacific warm pool (defined as SSTs \(>28.5\, ^\circ\mathrm{C}\)) that brackets the maritime continent of Indonesia, thereby preventing a shift of convection to the central Pacific Ocean in response to weak warm El Niño SST anomalies.

Lower-frequency variability may have played a role as well, as the background state of the tropical Indo-Pacific has been changing on decadal timescales. For instance, the Pacific Decadal Oscillation, or equivalently the Interdecadal Pacific Oscillation, had been in a cold phase since 1999, during which trade winds were of unprecedented strength and baseline SSTs were lower in the eastern and central equatorial Pacific. These cold background conditions, linked to the recent hiatus in global warming, could likewise have inhibited the migration of deep convection from the western to central Pacific in 2014. In addition, over the past 50 years the Indo-Pacific warm pool has been warming and expanding in response to greenhouse gas forcing. These warming trends may have contributed to the competition between Indo-Pacific warm pool SSTs and central Pacific SSTs in affecting the location of deep convection along the Equator in 2014. Interestingly, a budding El Niño died prematurely just two years earlier in 2012, perhaps under the influence of these same changing background conditions.

We have come to appreciate that there is a lot of natural variability in the ENSO cycle and that individual events can exhibit a wide diversity of behaviors. Evolution of conditions in 2014–2015 emphasizes just how complex ENSO dynamics can be. Climate change may plausibly have played a role via warming trends in the Indo-Pacific warm pool, but it will be a challenge to convincingly tease out its possible influence from a broad spectrum of natural variations. Seeking answers for why an erstwhile major El Niño unexpectedly went into hiding in 2014 only two years earlier in 2012, perhaps under the influence of these same changing background conditions.

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